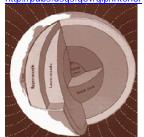
A discussion of the earth's interior is the basic prerequisite for understanding

Earth's dynamism partly because of the role of planetary dynamics in creating and sustaining the planetary processes and partly because the tools that inform us of the planetary interior are also the ones we need in our understanding of the surface effects.

Access the USGS publication "Interior of the Earth" at http://pubs.usgs.gov/gip/interior



		Whole Earth	Earth's Crust
Oxygen	0	29.8%	46.6%
Silicon	Si	15.6%	27.7%
Magnesium	Mg	13.9%	2.1%
Iron	Fe	33.3%	5.0%
Aluminum	AI	1.5%	8.1%
Calcium	Ca	1.8%	3.6%
Nickel	Ni	2.0%	
Sodium	Na	0.2%	2.8%
Potassium	K		2.6%
Others		1.9%	1.5%

Earth is a multi-layered body

- Crust is the earth's thin (0-70 Km) outer skin, averaging ~30 Km beneath the continents and ~15 Km beneath the oceans.
- Mantle is the earth's ~2,900 Km thick and rocky outer shell that underlies the crust.
- Core is the earth's ~3,500 Km thick metallic interior, comprising the (a) solid inner core (1,250 Km radius), and (b) liquid outer core (2,250 Km thick).

The undulations of the geoid, or the equipotential surface, reveal inhomogeneous mass

distribution. For instance, notice in this picture

This is NASA's picture of the earth geoid

(http://ekman.unh.edu/course/intro

the region immediately south of Africa.

(Inner core and crust have similar volumes of $\sim 8.2 \times 10^9$ Km³)

Of the 9 elements that dominate the Earth's chemical composition,

- crust carries most of the Earth's Si, O, Al, Ca and Na;
- mantle is, in effect, a compositional replica of the whole earth; whereas
- most of the earth's Fe, Mg and Ni occur in the core.

Information on the Earth's internal structure comes from (a) gravity, (b) seismic, and (c) geomàgnetic studiés.

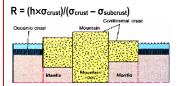
The Gravity Picture

Because of earth's equatorial bulge and polar flattening, gravitational acceleration on the surface increases from equator to the poles.

Isostasy

Earth's Interior

G.B. Airy's 1855 model of "isostacy", illustrated here, assumed that the low density crust floats over a denser sub-stratum. The root 'R' of crust beneath a mountain (or its antiroot beneath an ocean) is:



Here h is height (or depth) relative to mean sea level while σ_{crust} and $\sigma_{subcrust}$ are densities of crust and subcrust respectively.

- The whole earth density (~5.5 g/cm³) is about twice the average crustal density (~2.7 g/cm³): Clearly, density increases with depth.
- The continental crust is lighter
- and thicker than the oceanic crust, as "mountains have their own roots" (i.e., isostasy).

The Seismic Evidence

Seismic waves are of two types: the surface waves and the body waves (the P and S waves).

- Of these, P-wave velocity in the crust averages ~6 Km/s, S-wave ~4 Km/s.
- Earthquake focal depths are usually <250 Km.
- **The "shadow zone":** No direct P-waves from an earthquake. arrive between 103° and 142° from the epicenter, whereas no direct S-waves from an earthquake are seen beyond 103° from the epicenter. Since S-waves do not traverse a fluid layer, where the P-waves slow down, this suggests that the outer core is a fluid layer.

The Geomagnetic Field

- Earth has a magnetic field that behaves as if there is a bar magnet inside the earth and along the spin axis. This time- averaged geomagnetic field is a geocentric-axial-dipole.
- Crustal magnetization is too weak to produce this magnetic field and the sub-crustal region is too hot to be magnetic. The magnetohydrodynamics of the fluid and metallic, and therefore electrically

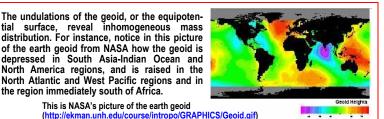


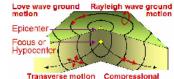
conducting, Outer Core offers the most acceptable self-sustaining and regenerating mechanism that we need to explain the origin of the geomagnetic field.

Why is the Outer Core Fluid?

Apparently because temperatures there exceed the melting points of core's constituents.

To learn about the geomagnetic field, visit the URL: http://www.ngdc.noaa.gov/seg/potfld/geomag.shtml





motion of P-waves of S-waves

